Abstract

Beach sands from the Rosa Marina locality (Adriatic coast, southern Italy) were analysed mainly microscopically in order to trace the source areas of their lithoclastic and bioclastic components. The main cropping out sedimentary units were also studied with the objective to identify the potential source areas of lithoclasts. This allowed to establish how the various rock units contribute to the formation of beach sands. The analysis of the bioclastic components allows to estimate the actual role of organisms regarding the supply of this material to the beach. Identification of taxa that are present in the beach sands as shell fragments or other remains was carried out at the genus or family level. Ecological investigation of the same beach and the recognition of sub-environments (mainly distinguished on the basis of the nature of the substrate and of the water depth) was the key topic that allowed to establish the actual source areas of bioclasts in the Rosa Marina beach sands. The sedimentological analysis (including a physical study of the beach and the calculation of some statistical parameters concerning the grain-size curves) shows that the Rosa Marina beach is nowadays subject to erosion.

Keywords: source-area reconstruction, grain-size analysis, mineralogical composition, image analysis, coastal erosion
1. Introduction

Beaches attract much interest from earth scientists nowadays (Schwartz, 2005), in particular from sedimentologists (Greenwood & Davies, 1984) and geomorphologists (Bird, 2008). Understanding the mechanisms which regulate the evolution of the beaches is not a simple scientific exercise, but understanding of these mechanisms may have an important economic and social impact; numerous scientific articles and books therefore focus on a variety of geological features concerning beaches. These features include shoreline dynamics (Ingle, 1966; Fredsøe & Deigaard, 1994; Anthony, 2009), coastal erosion (Charlier & De Meyer, 1998; Uda, 2010; Van Rijn, 2011), management and monitoring of coastal areas (NRC, 1989; Kay & Alder, 2002) and beach nourishment (Finkl, 1981; NRC, 1989; Nordstrom, 2005).

Beaches form an environment where ecological and biological factors tend to play a pivotal role with regard to the physical evolution of the system; nonetheless only few studies describing the interactions between the biological processes and the sedimentary dynamics in coastal areas have been published so far (NRC, 1994a,b).

In shallow-marine environments, sediment may result from the redistribution (due to waves, tides and currents) of the material supplied by rivers and/or eroded from rocks in the coastal area (both types form terrigenous clastic material) or from the production of clastic particles in the marine environment by biotic and/or abiotic processes (autochthonous clastic material). Sediments generated by biotic processes are most commonly carbonates; beach sediments therefore tend to contain a variable percentage of carbonates, derived from bioclasts (i.e., shells or fragments of marine organisms).

The study of the complex interactions between physical and biological processes provides key information for monitoring, protection and restoration of coastal areas. For example, beach nourishment by supply of sand that replaces eroded sediment (Chiocci & La Monica, 1999; Van der Salm & Unal, 2003; Nicoletti et al., 2006; APAT-ICRAM, 2007; Anfuso et al., 2011) needs: - a physical characterisation of the site of interest; - a chemical characterisation of the sediments to be supplied (with particular attention for their organic and inorganic pollution); - detailed biological and ecological analyses on the benthic populations, the presence of seagrasses, nursery areas, as well as attention to the impact of beach nourishment on the biotope (Coloso et al., 2007). Furthermore, the suitability of sand (coastal or relict) for beach nourishment ultimately depends on its ‘compatibility’ with the sands of the coast to be replenished; the definition of ‘compatibility’ requires a quantitative evaluation of the textural, petrographical and mineralogical parameters of the main (both lithoclastic and bioclastic) sediment constituents (colour, sieve size intervals of 1/2 φ, main mineralogical characteristics, etc.).

The present contribution describes a multidisciplinary approach in the study of a beach. It focuses on the interaction between the organisms living in the depths of the coastal areas and the physical processes connected to sedimentation. Both aspects are closely related, especially with regard to sediment production: the relationship between lithoclastic and bioclastic sediment in the Rosa Marina coastal area, along the Adriatic sector of the Apulian region is shown. A methodology for beach-sediment analysis is presented, aimed at both textural/petrographical characterisation of the sands and at the definition of the bioclast content as related to benthic populations and their relationships with the sandy or rocky substrate.

2. Geological setting

The Rosa Marina beach (N40°50’, E17°50’) is located north of Brindisi, along the southeastern Adriatic coast of Italy (Fig. 1). It is characterised by a great touristic and economic pressure, housing two private tourist-residential consortia (for a total of over 1,000 villas), a hotel-resort, swimming pools, etc. It is part of the Regional Natural Park of the Coastal Dunes between ‘Torre Canne’ and ‘Torre San Leonardo’ (established by Regional Law 31, dated 26–10–2006). This coastal area includes small catchment areas through which ephemeral streams run (‘Pilone Vallone’ and ‘Rosa Marina Lama’: Fig. 1) that are capable to carry only moderate amounts of sediments to the sea, usually during intense weather events. This contributes to the building of a coastal wedge in a microtidal setting where the littoral dynamics are wave-dominated. The wave conditions, deduced from the 1968–2008 data of a wave buoy at Monopoli (N40°58’30.0”, E17°22’36.1”, less than 25 km from the beach under study) indicate that the prevailing direction of sea storms is from the north-west. The sediment transport by traction along the coast should, therefore, mainly occur in a NW-SE direction (Mastronuzzi et al., 2002).

This Apulian coastal sector often shows a linear trend, with low-elevation, active sandy coastal dunes and well-developed backshore marsh areas (south of Torre Canne). From Pilone to the southeastern Monticelli (an area including Rosa Marina),
the coast becomes more irregular with rocky areas representing local headlands. Some sandy pocket beaches (narrow beaches between two headlands) with inactive and/or erosion-affected rows of dunes are preserved.

According to the Regional Plan of Coasts (Puglia Region, 2012), the coastal area under study is characterised by intense urbanization; direct interventions on the coast concerned, however, only a transverse barrier at the mouth of the ‘Rosa Marina Lama’ stream and a short breakwater pier in the same area. Phenomena of coastal erosion have been reported in Annese et al. (2003) for the Torre Canne area (about 7 km north of Rosa Marina), whereas the shorelines were stable in the Pilone area (Fig. 1) (Puglia Region, 2006).

The geological setting (Fig. 2) of this area has been described in detail (Ciaranfi et al., 1988; Martonuzzi et al., 2001). Nevertheless, a preliminarily, short geological and sedimentological survey was carried out in the study area. It was focused on the recognition of the main outcropping sedimentary units, the location of their stratigraphic boundaries with reference to the average sea level, the lateral variations in their facies and/or thickness, and their state regarding erosion (Fig. 2). All the outcropping sedimentary units have been sampled for petrographical analysis. The Calcarenite di Gravina Formation (Pliocene – Early Pleistocene) is the oldest sedimentary unit cropping out in this area. It represents the substratum of a succession of marine terraced units deposited during the Middle Pleistocene to Holocene (Ciaranfi et al., 1988). The Calcarenite di Gravina Fm. is present almost continuously along the entire coast of the study area, though mostly below sea level, where it is abraded by storm waves (GRA, Fig. 2A). In this area, the Calcarenite di Gravina Fm. is made up of thick beds of medium- and coarse-grained calcarenite, with intense bioturbation and a high content of bioclasts (mostly red algae) (Moretti et al., 2011).

On top of the Calcarenite di Gravina Fm., a thin red soil unit has been recognised over a continental erosional surface, approximately at sea level (TR, Figs 2A, B): it passes upwards and laterally into marly limestones (C1). Unit C1 is made up of parallel-laminated, fine-grained limestones with abundant ostracods and with rare clay chips of red soils. On top of the TR and C1 units, a coarse-grained calcarenite unit crops out, a few centimetres above sea level (C2, Figs 2A, B). It is made up of parallel-laminated calcarenites to calcirudites with a high bioclast content (mainly fragments of bivalves and gastropods). Laterally, close to the ‘Rosa Marina Lama’ stream, unit C2 passes into a calcirudite (unit CR) with some gastropod remains and abundant pebbles of micritic limestone and calcarenite (Figs 2A, C).

Between 10 and 30 cm above sea level, overlying older units, an arenaceous unit (E1) crops out (Figs 2A, C). It is made up of well-sorted sand/sandstone alternations with high-angle cross-lamination; the sands have a mixed composition, with quartz and carbonates in almost similar percentages. The top-
most outcropping unit is sandy; its base is situated 40 cm to 2 m above sea level. It is a subrecent aeolian unit (E2), representing a coastal dune that is no longer active nowadays, but instead exposed to strong erosion (Figs 2A, C).

The succession described above represents the result of interaction between the slow uplift of the Apulian Foreland and sea-level changes. The informal units, TR, C1, C2 and E1, are considered to have been deposited on top of the Gravina Calcarenite, during a transgressive/regressive cycle, recorded by units of coastal lagoon/backshore sediments (TR and C1), passing upwards into shoreface transgressive deposits (C2, laterally fed by the terrigenous carbonates supplied by the ‘Rosa Marina Lama’ stream, CR). The aeolian unit (E1) represents the regressive part of the succession: it is considered to have been deposited in this area, as in many other coastal areas of the Apulian Foreland, during the middle Holocene (about 6,000 years ago: Mastro-

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**Fig. 2.** Geology of the Rosa Marina area.

A – Schematic stratigraphy of the Rosa Marina area. See the text for the complete names and ages of the sedimentary units; B – Stratigraphical contact (in white) between the red soil unit (TR, at the bottom) and the overlying lagoonal calcarenite (C1). The yellow line indicates the contact between the lagoonal unit (C1) and the overlying transgressive calcarenite (C2); C – Macroscopic features of the of the calcirudite unit (CR), which passes laterally to the C2 calcarenites; D – Contact between the middle Holocene aeolian unit (E1) and the overlying younger dune (E2). Note that, on the foreground, unit E1 undergoes considerable erosion, with localized falls and block rotation; in the background, unit E2 is directly exposed to wave erosion.
nuzzi et al., 2001; Mastronuzzi & Sansò, 2002). The more recent coastal dune E2, however, was presumably deposited during the late Holocene (Mastronuzzi et al., 2001).

3. Methodology

The sands of the present-day beach were sampled in both emerged and submerged areas (Fig. 3): (1) in the shoreface, along a transect perpendicular to the coast, from the shoreline to a depth of 6 m (the local storm-wave base), and (2) in the backshore and foreshore, where sand samples were taken every 5 m from the shoreline, taking care to sample both the ordinary and winter berms (in lateral and less frequented areas) until the base of the E2 aeolian sands. A biological survey was carried out through diving at the investigated depth (up to -6 m) and through sampling of organisms from both the sandy and the hard substratum.

The samples for the sedimentological analyses were collected by driving a cylinder sampler, which was tightly closed in order to avoid loss of finer sediments. In the laboratory, the samples were washed with distilled water, dried and weighed. They were processed with hydrogen peroxide and subsequently passed through a 0.063 mm sieve in order to determine the percentage of organic matter and fine sediment (both present in negligible percentages). The results were processed with the Gradistat v8 software (Blott & Pye, 2001), which yielded distribution histograms, cumulative curves, and the automatic evaluation of the following textural parameters: median grain size ($D_{50}$), sorting (sg), skewness (Sk) and kurtosis (Kg). In order to obtain ecological data, all bioclasts (shells or fragments of shells) were separated from the remainder of the sedimentary particles and placed in special Petri dishes using a binocular optical microscope. All sedimentary units cropping out in the study area were also sampled and studied petrographically for comparison in order to obtain more information on the role of erosion of the local substratum as a source for the present-day beach sands. First, the qualitative composition of beach sands was obtained with the help of a binocular optical microscope. More quantitative analyses of the present-day beach sands were carried out after the epoxy resin impregnation of the present-day beach sands and obtaining slabs for thin sections; five thin sections (covering both emerged and submerged sub-environments) were analysed, and high-resolution microscope photos were taken using low magnifications (x1 and x2). The raster images that were thus obtained were imported into image-analysis freeware (ImageJ, version 1.49) with the help of which it was possible to recognise, select and draw individual grains and to assign them to different compositional classes.

Fig. 3. Locations of the sampling stations. In the backshore (Ba), samples 1–4 are located at 20, 15, 10 and 5 m from the swash zone, respectively. Sample ‘dune’ was collected at the base of the E1 dune unit, whereas samples Bo and Bt (ordinary and winter berm) were collected laterally (to the south-east of the transect) in an area of the beach where such morpho-sedimentary steps were still recognisable.
4. Grain-size analysis

The present-day beach sediments are, from a granulometric point of view, medium- to coarse-grained sands (Table 1 and Fig. 4).

Typically, relatively high $D_{50}$ values were found for the berms; the values decrease in the shoreface environments with increasing water depth (Fig. 5). No bars or sediment accumulation areas have been detected in the submerged sectors (Fig. 5). The sands are mostly well-sorted; sorting decreases with water depth. The skewness is equal to zero on the shoreline (Fig. 3) and negative on the backshore profile, resulting in a tail of coarse material (mean left of the median); at the same points, Kg (i.e., the ratio between the width of the central part of the diagram and that of the tail) shows high values (leptokurtic type), except for sample Bt (curve of the platykurtic type).

The grain-size parameters can be used also for obtaining (essentially qualitative) information on the evolution of a beach with particular reference to the susceptibility to erosion (Dal Cin, 1969).

Table 1. Grain-size parameters determined for the samples from the beach sands.

<table>
<thead>
<tr>
<th>Sample</th>
<th>$D_{50}$ (μm)</th>
<th>Sorting ($σφ$)</th>
<th>Description</th>
<th>Skewness $(Skφ)$</th>
<th>Shape</th>
<th>Kurtosis $(Kg)$</th>
<th>Shape</th>
</tr>
</thead>
<tbody>
<tr>
<td>dune (E2)</td>
<td>399.695</td>
<td>0.470</td>
<td>medium sand, well sorted</td>
<td>0.568</td>
<td>tail to the fine fraction</td>
<td>1.676</td>
<td>strongly leptokurtic</td>
</tr>
<tr>
<td>1</td>
<td>417.618</td>
<td>0.365</td>
<td>medium sand, well sorted</td>
<td>0.051</td>
<td>symmetrical</td>
<td>2.482</td>
<td>strongly leptokurtic</td>
</tr>
<tr>
<td>winter berm (Bt)</td>
<td>782.277</td>
<td>0.498</td>
<td>coarse sand, well sorted</td>
<td>0.562</td>
<td>tail to the fine fraction</td>
<td>0.631</td>
<td>strongly leptokurtic</td>
</tr>
<tr>
<td>2</td>
<td>422.702</td>
<td>0.416</td>
<td>medium sand, well sorted</td>
<td>-0.006</td>
<td>symmetrical</td>
<td>2.636</td>
<td>strongly leptokurtic</td>
</tr>
<tr>
<td>3</td>
<td>433.633</td>
<td>0.577</td>
<td>medium sand, moderately well sorted</td>
<td>-0.299</td>
<td>tail to the coarse fraction</td>
<td>2.446</td>
<td>strongly leptokurtic</td>
</tr>
<tr>
<td>4</td>
<td>429.620</td>
<td>0.302</td>
<td>medium sand, very well sorted</td>
<td>-0.292</td>
<td>tail to the coarse fraction</td>
<td>1.77</td>
<td>strongly leptokurtic</td>
</tr>
<tr>
<td>ordinary berm (Bo)</td>
<td>800.881</td>
<td>0.467</td>
<td>coarse sand, well sorted</td>
<td>0.579</td>
<td>tail to the fine fraction</td>
<td>1.737</td>
<td>strongly leptokurtic</td>
</tr>
<tr>
<td>backshore (Ba)</td>
<td>838.784</td>
<td>0.155</td>
<td>coarse sand, very well sorted</td>
<td>0</td>
<td>symmetrical</td>
<td>0.738</td>
<td>platykurtic</td>
</tr>
<tr>
<td>-1 m</td>
<td>392.986</td>
<td>0.576</td>
<td>medium sand, moderately to well sorted</td>
<td>0.319</td>
<td>tail to the very fine fraction</td>
<td>0.836</td>
<td>platykurtic</td>
</tr>
<tr>
<td>-3 m</td>
<td>397.568</td>
<td>0.825</td>
<td>medium sand, moderately to well sorted</td>
<td>-0.040</td>
<td>symmetrical</td>
<td>0.847</td>
<td>platykurtic</td>
</tr>
<tr>
<td>-6 m</td>
<td>249.487</td>
<td>0.662</td>
<td>medium sand, moderately to well sorted</td>
<td>-0.558</td>
<td>tail to the very coarse fraction</td>
<td>0.850</td>
<td>platykurtic</td>
</tr>
</tbody>
</table>

Table 2. Grain-size parameters determined for the adjacent Pilone beach.

<table>
<thead>
<tr>
<th>Sample</th>
<th>Sampling depth</th>
<th>$D_{50}$ ($φ$)</th>
<th>Sorting ($σφ$)</th>
<th>Description</th>
<th>Skewness $(Skφ)$</th>
<th>Shape</th>
<th>Kurtosis $(Kg)$</th>
<th>Shape</th>
</tr>
</thead>
<tbody>
<tr>
<td>2006 3_15_7</td>
<td>+1.39</td>
<td>1.56</td>
<td>0.39</td>
<td>medium sand, well sorted</td>
<td>0</td>
<td>symmetrical</td>
<td>0.93</td>
<td>mesokurtic</td>
</tr>
<tr>
<td>2006 3_15_8</td>
<td>+0.58</td>
<td>1.85</td>
<td>0.48</td>
<td>medium sand, well sorted</td>
<td>0.03</td>
<td>symmetrical</td>
<td>1.04</td>
<td>mesokurtic</td>
</tr>
<tr>
<td>2006 3_15_15</td>
<td>-2.90</td>
<td>2.08</td>
<td>0.43</td>
<td>fine sand, well sorted</td>
<td>-0.07</td>
<td>symmetrical</td>
<td>1.33</td>
<td>leptokurtic</td>
</tr>
<tr>
<td>2006 3_15_16</td>
<td>-4.58</td>
<td>2.27</td>
<td>0.32</td>
<td>fine sand, very well sorted</td>
<td>0.09</td>
<td>symmetrical</td>
<td>1.54</td>
<td>strongly leptokurtic</td>
</tr>
</tbody>
</table>
Samples from the beach of Rosa Marina commonly show values of the various textural parameters which indicate an erosional-regressive evolutionary trend: (1) sorting is clear in the backshore-foreshore-shoreface sectors; (2) Sk is mainly negative (it is a clear record of severe storm-wave hydrodynamics); (3) the distribution curves are often leptokurtic or even very leptokurtic, with tails shifted mostly toward the coarse-grained material.

Of course, more reliable data might have been obtained if samples had been collected repeatedly throughout a year, and preferably during several years. Reference data with which our data can be compared are available, however, for adjacent areas (Pilone beach resort: Fig. 1); data are taken from the Puglia Region (2006): see Table 2. The Pilone beach is located in a stable (or relatively less erosional) coastal sector showing sub-symmetric distributions (Sk about 0) and mesokurtic to leptokurtic curves (Table 2). These data confirm the erosional trend established for the Rosa Marina area.

5. Petrographical analysis

5.1. The beach sands

The petrography of the beach sands was analysed for different size fractions in order to reveal the percentages of bioclastic particles and any variations in the composition and/or concentration of specific minerals in the various granulometric classes. First the bioclast content was determined quantitatively (the quality is discussed in Section 5.2). The bioclasts were weighted and their total fraction was calculated as the percentage of the total sample weight (Table 3).

The bioclast percentage by volume was calculated using a simple conversion. Siliciclastic particles and carbonate particles have more or less the same density: the quartz grains have a density of 2.66 g∙cm$^{-3}$ whereas that of the carbonates is about 2.70 g∙cm$^{-3}$; Bioclastic carbonates can, however, have different densities (Schlager, 2005) according to their composition (2.94 g∙cm$^{-3}$ for aragonite; 2.72 g∙cm$^{-3}$ for calcite; 2.89 g∙cm$^{-3}$ for dolomite),
and the structure of the shells and other fragments of marine organisms usually have a varying but relatively high porosity (intraparticle porosity of Choquette & Pray, 1970), resulting in a total bulk volume that varies from a maximum value of 2.7 g cm$^{-3}$ (porosity = 0) to a minimum of 2.0 g cm$^{-3}$ (Jackson & Richardson, 2007). To obtain the maximum volume percentage of the bioclasts (the maximum difference from the percentages calculated considering their weight: Table 3), we considered the bulk density of the more porous skeletal material (the "unfilled shells" of Choquette & Pray, 1970).

The bioclast content in the backshore and foreshore is low, but it increases linearly from the dune to the shoreline, where it becomes predominant (Table 3); in the shoreface it increases with water depth. The bioclast content varies in the different grain-size fractions, but the shells and fragments of shells are most characteristic of the coarse fractions (>500 μm).

The lithoclasts of the beach sands (Fig. 6) consist mainly of carbonates, quartz and other minerals that are present in negligible percentages (such as pyroxene and feldspar); there are also rare fragments of siliciclastic rocks and of anthropogenic material (Fig. 6B, C).

The carbonate lithoclasts predominate in the terrigenous portion of the sands. They are monomineralic fragments of older rocks (micritic limestones and rarely calcarenites) and polymineralic lithoclasts of calcite together with other minerals. The second component in the beach sands is quartz. It is almost exclusively crystalline quartz and only rarely chert or chalcedony. The quartz grains always are well rounded (Fig. 7B). Furthermore, some rare feldspars are present in the form of potassium feldspar or plagioclase. The potassium feldspar grains (mainly microcline) contain inclusions (zircon and plagioclase: Fig. 7C) and show alteration rims in the form of clay minerals (Fig. 7D). The dark-coloured minerals consist mainly of pyroxene (Fig. 7E), which shows up light green under crossed nichols and yellowish under parallel nichols; they often are well rounded and they may contain inclusions of volcanic glass (Fig. 7E). The quantitative data on the composition of these sands come from the image analysis (Fig. 8): three main compositional classes of sands (carbonates, quartz and other minerals)

### Table 3. Bioclast/lithoclast percentages (in weight and volume) in the beach sand.

<table>
<thead>
<tr>
<th>Sample</th>
<th>Bioclasts (% weight)</th>
<th>Lithoclasts (% weight)</th>
<th>Bioclasts (% V$_{max}$)</th>
<th>Lithoclasts (% V$_{min}$)</th>
</tr>
</thead>
<tbody>
<tr>
<td>dune (E2)</td>
<td>0.74</td>
<td>99.26</td>
<td>1.00</td>
<td>99.00</td>
</tr>
<tr>
<td>1 (20 m from the shoreline)</td>
<td>1.98</td>
<td>98.02</td>
<td>2.67</td>
<td>97.33</td>
</tr>
<tr>
<td>2 (15 m from the shoreline)</td>
<td>4.71</td>
<td>95.29</td>
<td>6.35</td>
<td>93.65</td>
</tr>
<tr>
<td>3 (10 m from the shoreline)</td>
<td>3.05</td>
<td>96.95</td>
<td>4.11</td>
<td>95.89</td>
</tr>
<tr>
<td>4 (5 m from the shoreline)</td>
<td>3.04</td>
<td>96.96</td>
<td>4.10</td>
<td>95.90</td>
</tr>
<tr>
<td>ordinary berm (Bo)</td>
<td>16.84</td>
<td>83.16</td>
<td>22.73</td>
<td>77.27</td>
</tr>
<tr>
<td>swash zone (Ba)</td>
<td>54.58</td>
<td>45.42</td>
<td>73.68</td>
<td>26.32</td>
</tr>
<tr>
<td>-1 m</td>
<td>1.06</td>
<td>98.94</td>
<td>1.43</td>
<td>98.57</td>
</tr>
<tr>
<td>-3 m</td>
<td>2.42</td>
<td>97.58</td>
<td>3.27</td>
<td>96.73</td>
</tr>
<tr>
<td>-6 m</td>
<td>11.36</td>
<td>88.64</td>
<td>15.34</td>
<td>84.66</td>
</tr>
</tbody>
</table>

Fig. 6. Petrographical features of the Rosa Marina sands observed with an optical binocular microscope.  
A – Grains with a diameter of >500 μm: limestone lithoclasts and several rounded quartz grains; B – Piece of polished metal from some jewelry (earring, necklace?); C – Large fragment of transparent glass (from a bottle).
Table 4. Composition (in percentages) of the Rosa Marina sands calculated in 10 different photos. The mean values can be considered as representative (note that values do not differ significantly) and indicate relative frequencies of carbonates as 62%, of quartz as 34%, and of other minerals as 4%. The ‘other’ class contains feldspar, pyroxene and other minerals that are present in negligible percentages.

<table>
<thead>
<tr>
<th></th>
<th>Carbonate (%)</th>
<th>Quartz (%)</th>
<th>Other (%)</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>59.33</td>
<td>36.42</td>
<td>4.25</td>
</tr>
<tr>
<td>2</td>
<td>62.06</td>
<td>32.73</td>
<td>5.21</td>
</tr>
<tr>
<td>3</td>
<td>63.24</td>
<td>33.4</td>
<td>3.36</td>
</tr>
<tr>
<td>4</td>
<td>63.43</td>
<td>35.35</td>
<td>1.22</td>
</tr>
<tr>
<td>5</td>
<td>62.91</td>
<td>32.21</td>
<td>4.88</td>
</tr>
<tr>
<td>6</td>
<td>58.05</td>
<td>37.34</td>
<td>4.61</td>
</tr>
<tr>
<td>7</td>
<td>62.43</td>
<td>33.26</td>
<td>4.31</td>
</tr>
<tr>
<td>8</td>
<td>61.78</td>
<td>34.86</td>
<td>3.36</td>
</tr>
<tr>
<td>9</td>
<td>60.73</td>
<td>35.29</td>
<td>3.98</td>
</tr>
<tr>
<td>10</td>
<td>64.67</td>
<td>33.91</td>
<td>1.42</td>
</tr>
<tr>
<td>Mean value</td>
<td>61.86%</td>
<td>34.48%</td>
<td>3.66%</td>
</tr>
</tbody>
</table>
Two samples, viz. Bo (ordinary berm) and the sample collected at a water depth of 6 m, are close to the field of hybrid sand. The sample from the shoreline (Ba) has a bioclast content that classifies it as a hybrid sand if we consider the weight percentage, and as a bioclastic hybrid sand on the basis of the volume percentage.

5.2. The older sedimentary units

The relatively thick units along the investigated shoreline that are affected by erosion (Fig. 2) are formed by the Calcarenite di Gravina Fm. (GRA) and by the recent aeolian dunes (E1 = middle Holocene and E2 = late Holocene). The Calcarenite di Gravina Fm. consists locally of a massive bioclastic sand with a packstone texture (Fig. 10A); it is made up almost entirely of red algae (with rare lithoclasts, fragments of bivalves and benthic foraminifers). The middle Holocene aeolian unit (E1, Fig. 10B) has the same petrographic characteristics as the present-day sands except for the presence of a calcite cement. Carbonate clasts (which form over 60% of the rock), quartz and some pyroxene (with petrographic characteristics very similar to those of the present-day sands) are present.

The present-day sands thus clearly show a compositional affinity with the recent aeolian-dunes (E1 and E2) and contain also numerous clasts of fossil
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6. Ecology of the beach

The area of interest was also analysed from an ecological/biological point of view in order to obtain a clear picture of the distribution of the main biocenoses in both sandy areas and in the hard-rock substratum. Single bioclasts of organisms that contribute with their remains to the sands have been collected and identified (commonly at genus level). The distributions of the organisms on soft- and hard- bottom yields evidence regarding the origin of the bioclasts in terms of bathymetric and ecological zones.

6.1. The main biocenosis

The analysed area is situated between the supralittoral and the upper infralittoral zones (sensu Peres & Picard, 1964); it can be divided into (1) the backshore, corresponding to the supralittoral zone; (2) the foreshore, corresponding to the mesolittoral zone; and (3) the shoreface until the wave base (~6 m), corresponding to the upper infralittoral zone. In all parts of the study area, soft sediment and hard substratum alternate. In the supralittoral zone, the biocenosis of the supralittoral sands is characterised by the presence of bioclastic material deposited during severe storms (algae, seagrass, remains of terrestrial plants, remains of marine and terrestrial organisms). This biocenosis alternates with that of supralittoral rocks colonized by few organisms such as gastropods of the genus *Melaraphe*.

In the mesolittoral zone, the upper intertidal rock contains a biocenosis that is characterised by deposits of cyanobacteria, crustaceans, barnacles like *Chthamalus* and gastropods such as *Patella*. In the lower intertidal zone, vermetids occur, indicating a priority habitat for the marine conservation in the SPA/BIO Protocol (Specially Protected Areas and Biological Diversity in the Mediterranean) of the Barcelona Convention (1997; see also Relini & Giaccone, 2009). The facies is characterised by bioconstructions of the sessile gastropod *Dendropoma petraeum* that builds complexes which induce a large increase in animal biodiversity (especially annelids, molluscs, crustaceans, echinoderms, and small benthic fish) and vegetation (calcareous seaweed thallus, algal mats and leafy algae). This highly diversified habitat is particularly sensitive to oil pollution and surfactants as well as to mechanical destruction related with the harvesting of some bivalves of the genus *Lithophaga*.

In the infralittoral zone, the soft-sediment sea floor is locally characterised by a biocenosis of infralittoral algae. At intermediate depths (~3 m), the bedrock is only sparsely inhabited, probably as a consequence of wave-induced erosion. At a depth of ~6 m, the biocenosis is more diversified and characterised by encrusting and leafy algae, sponges (especially the photophilous species *Chondrilla nucula*, which covers large portions of the hard substratum:}

![Petrographical details in thin section.](image)
6.2. Identification of the bioclasts

The bioclastic content (Table 4) of the beach sands has been analysed in more detail: the remains of the organisms were identified and classified, as far as possible, from a taxonomical point of view. The bioclasts consist of fragments of (1) rhizopods, (2) mollusc shells, (3) thorns or fragments of echinoderm exoskeletons, (4) bryozoans, and (5) fragments coming from less frequent organisms (this fifth category, called “other”, includes the remains of algae, spicules of sponges, fragments of serpulid pipes, fragments of barnacles and other crustaceans).

The largest contribution comes from molluscs, in particular from gastropods and bivalves, although the relative percentages are highly variable in the analysed sub-environments (Fig. 12). The number of identified mollusc taxa is 55 (Table 5, Fig. 13); the other bioclasts come from 39 taxa of gastropods (37 identified at the genus level and second-level family) and 16 taxa of bivalves (15 identified at the genus level and 1 at the level of the family). The number of taxa (Fig. 14) tends to increase gradually from the dunes (only one taxon) to the shoreline (52 taxa), to decrease again with water depth; it finally increases again at a depth of –6 m (18 taxa).

The recognised mollusc taxa are not typical of a single area, but rather come from slightly different ecological and bathymetric sub-environments (Fig. 15). A conspicuous part of the molluscs (36.4% of the total) is typical of rocky bottoms; the more frequent taxa, found at three or more sampling stations, belong to the genera *Columbella* (Fig. 13D) and *Striarca* (family Chamidae: Fig. 13H). Another large fraction (23.6% of the total) is characteristic of sandy bottoms; the most frequent taxa (also found at three or more sampling stations) belong to the genera *Caecum, Nassarius, Venericardia*, and *Parvicardium*. Another fraction (10.9% of the total) consists of mol-
There are also seven taxa of gastropods (12.7% of the total) that prefer to live in the grasslands of seagrass or algae. The most frequent taxa (found at three or more sampling stations) are the genera *Tricolia* (Fig. 13B), *Jujubinus*, *Rissoa* and *Alvania*. One taxon, the genus *Bittium* (Fig. 13A), is practically ubiquitous and can live on shallow rocky bottoms covered with vegetation, in the grasslands of marine plants and also on soft substratum. In particular, shell fragments of *Bittium* have been found in all analysed samples. Only two taxa (3.6% of the total) are typical of slightly deeper environments; they are the genera *Clanculus* (Fig. 13C) and *Irus*.

Five taxa (9.1% of the total) are parasites of gastropods or of other organisms: the genus *Epitonium*

**Fig. 13.** Well preserved bioclasts within the Rosa Marina beach sands.  
A – *Bittium* is the only taxon that is present in all samples; B – Gastropods of the genus *Tricolia* are typical of seagrass; C – *Clanculus* is a gastropod typical of rocky environments, particularly at at depth of at least 10 m; D – Gastropods of the genus *Columbella* are characteristic of shallow rocky environments with a rich vegetation; E – *Fossarum* is a gastropod that commonly is found in a rocky mesolittoral zone; F – *Melaraphe* is a gastropod of rocky environments in the supralittoral zone; G – Gastropods of the genus *Hydrobia* live in brackish transitional areas; H – *Striarca* is a bivalve that is common as bioclasts in the Rosa Marina beach sands; it is typical of a rocky environment.

**Fig. 14.** Number of taxa in the various beach sub-environments.
is an anemones parasite, the Triphoridae family lives mainly on sponges, Melanella is typically an echinoderms parasite, the genus Odostomia is a parasite of worms and shellfish, and the genus Cerithiopsis feeds mainly on sponges that are its habitat. Finally, one taxon, the genus Hydrobia (Fig. 13G), is typical of brackish environments.

7. Discussion

Geomorphological and sedimentological analyses show that the Rosa Marina beach is now distinctly subjected to erosion; evidence consists of the local erosion of the coastal dunes, and the lack of detectable depositional bars in the shoreface. Moreover, the calculation of all statistical parameters regarding the grain sizes of the sediments at various places along and perpendicular to the coast indicate that the Rosa Marina area is being eroded, whereas adjacent coastal sectors seem to be more stable. The petrographical analyses of sand samples carried out with a binocular and in thin section (using also some image-analysis procedures) indicate regarding their modal composition that they are made up of calcium carbonates (about 62%), quartz (about 34%) and other minerals (K-feldspar, plagioclase, pyroxene, etc. constitute about 4%). A similar petrographical composition has been found for the rocks cropping out in the coastal area showing the importance of the cannibalization processes for the recent formation and evolution of coastal deposits in the Apulian Foreland, which is consistent with regional data
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(Tropeano et al., 2002; Tropeano & Spalluto, 2006; Gallicchio et al., 2014; Gioia et al., 2014). The quantitative evaluation of the bioclastic components, obtained by physical separation and by weighing the shells and shell fragments, shows that this fraction varies strongly: from a minimum of 1% to a maximum of over 50% by weight (this equals some 70% by volume). The distinction between bioclastic and lithoclastic particles indicates that the beach sands should be classified as carbonate extraclastic hybrid sands, except for one sample which can be classified as a hybrid sand or a bioclastic hybrid sand, depending on whether the weight or volume of the various components is considered.

The ecological/biological survey carried out in the various marine sub-environments made clear which are the source areas of the various kinds of bioclasts that are present in the sands. Most of the bioclasts are shells and fragments of bivalves and gastropods. Forty taxa (72.7% of the total) come from sandy and rocky submerged shallow environments, in particular from the biocenosis of infralittoral algae and from the facies with vermetids. A smaller number of taxa (16.4% of the total) comes from deeper zones, especially from seagrasses and from coralligenous complexes.

Finally, a group of mollusc shells belongs to parasites that live in various sub-environments; only one taxon is typical of brackish environments. It seems evident that there are multiple sources of bioclasts and that these sources are characterised by different ecological parameters (salinity, depth, light, type of substratum, presence of vegetation, presence of bioconstructors). Nevertheless, our study shows that the most important source areas are located within the shoreface and offshore transition environments (only 16.4% comes from deeper environments).

8. Conclusive remarks

We have analysed sands of a beach located along the Adriatic coast of the Apulian region at Rosa Marina (north of Brindisi) with the objective to characterise in detail the lithoclast and bioclastic components, so that the source of the beach sands could be traced. The beach was analysed for the purpose from both a physical point of view (geomorphology and sedimentology) and an ecological/biological point of view. This approach was considered the most appropriate because beaches are the result of a combination of several physical processes (erosion, transport and sedimentation) and biological agents that interact at various scales and at different times.

A first important conclusion is related with the comparative analyses carried out concerning the composition of the present-day beach sands and regarding the sedimentary units cropping out in the coastal area. We come to the conclusion that the lithoclast component of the present-day beach sands is controlled by storm-wave erosion of the rocky substratum, without a necessary contribution from longshore sediment transport. Furthermore, the sedimentary units of the recent coastal dunes have a composition that is highly similar to the present-day beach sands; this conclusion seems only logical considering the unchanged or similar palaeogeographical situation during the late Pleistocene-Holocene.

Another important conclusion regarding the coastal erosion can be drawn from the bioclastic content: also in beaches with an insignificant bioclastic content (in one sample it was close to only 1%), such as in the case of the study area, a considerable coastal retreat can occur in only a few years by the loss or by a drastic decrease of the bioclastic input. This conclusion indicates how important a detailed study is of the bioclastic components of the sands by identification of the taxa that supply sedimentary material to the beach. The natural bioclastic supply to the beach depends on the health state of these environments and on the possibility to host, to feed, to provide shelter and to be a breeding site for taxa that, one day, will become bioclasts.

The results of this study also indicate how effective and less expensive procedures can be developed to prevent or minimize coastal erosion. In this context we think it of potentially great social importance to emphasise the enormous potential of geological/biological multidisciplinary approaches in the study of beaches: most important are (1) the characterisation of the physical/biological/ecological environment, (2) the analysis and monitoring of
the phenomena involved in coastal retreat, (3) the investigation and characterisation of materials that can be useful for beach nourishment, and (4) the analysis of ecological impacts in the widest sense at every stage of beach nourishment.

References


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